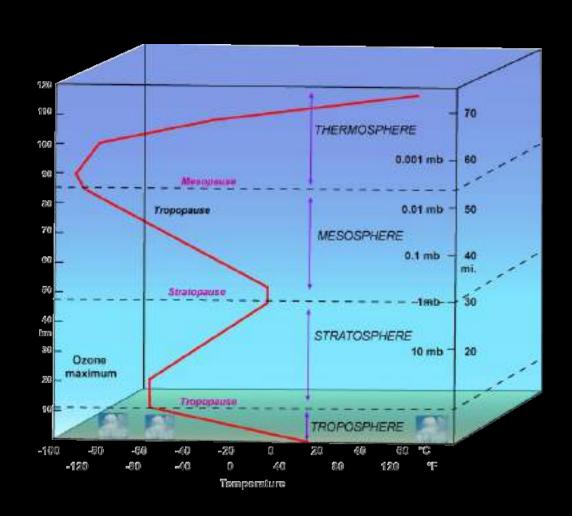
Climatology Tapan Debnath

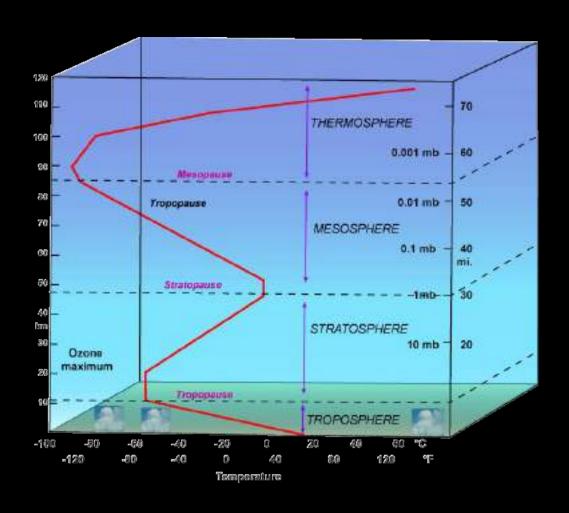
Topic

- 1. Structure of Atmosphere
- 2. Laps rate
- 3. Inversion of Temperature
- 4. Stability and Instability
- 5. Humidity
- 6. Precipitation
- 7. Pressure System
- 8. Jet stream and Indian Monsoon
- 9. Airmass

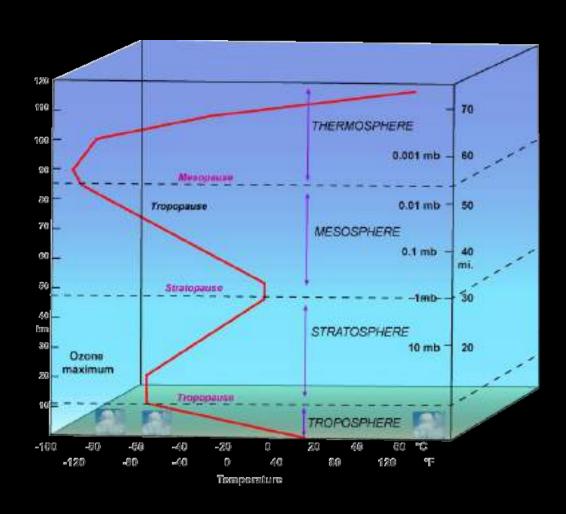


Troposphere

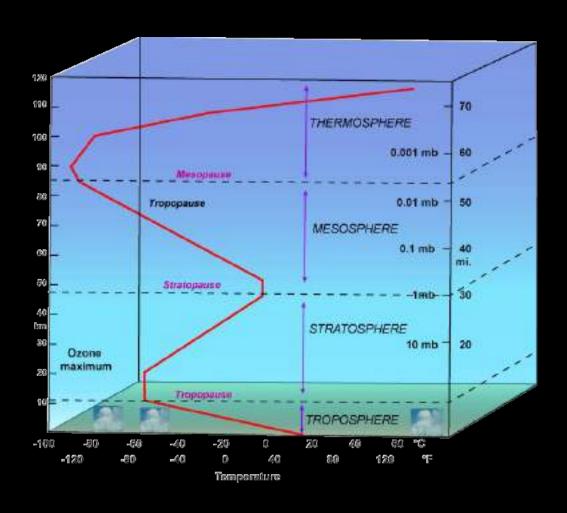
- Temperature decreases with increasing height
- GHG gas absorb long wave terrestrial radiation – transparent to insolation (short wave radiation)
- Heated from terrestrial radiation from earth surface



- Stratosphere
- Temperature increases with increasing height
- Ozone layer absorbs UV rays from Insolation



- Mesosphere
- Temperature decreases with increasing height
- Absence of GHG gas
- Going away from Ozone layer

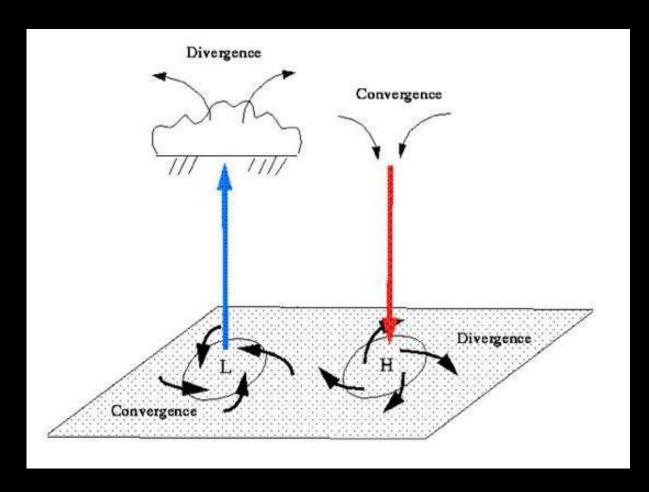


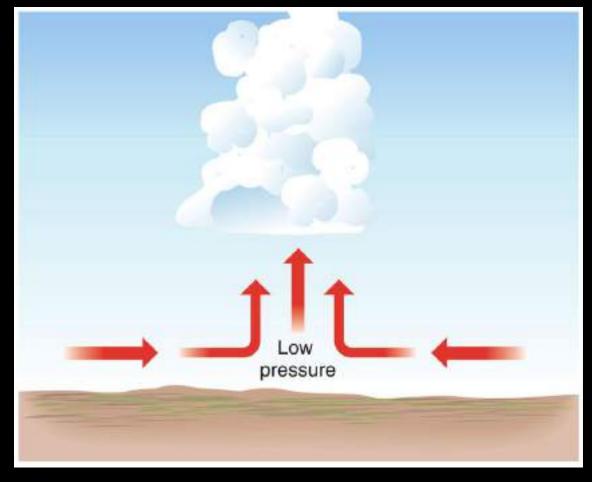
- Thermosphere
- Temperature increases with increasing height
- Gases are in ionic state–extremely hot
- Ionosphere -Imp for radio communication

Discuss

Basic air movement

Adiabatic lapse rate



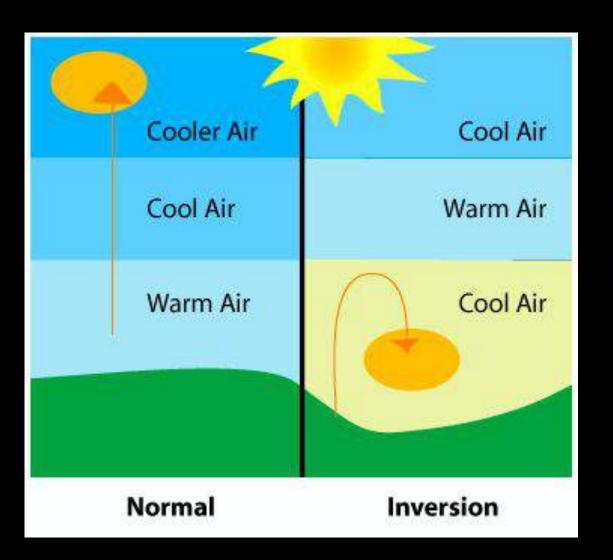


Change of Temperature

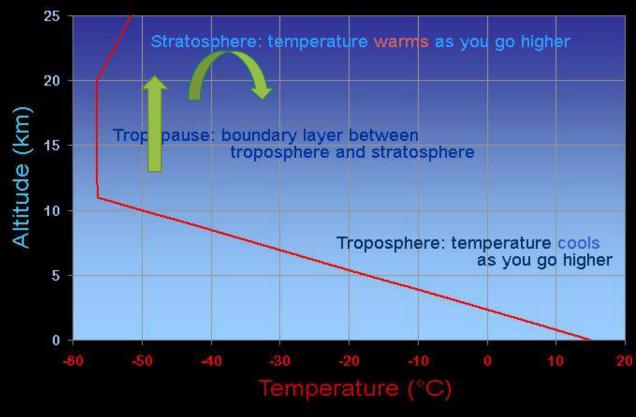
1. Diabetic Processes

2. Adiabatic Processes

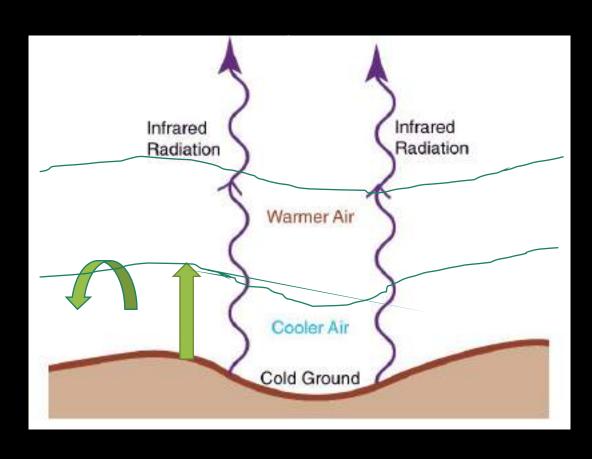
Temperature Inversion

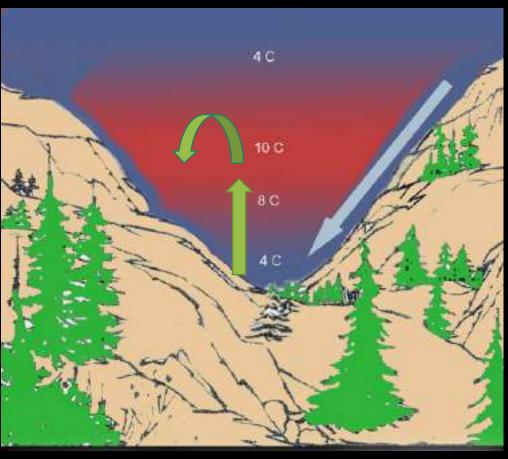


 Ex. Of Temp Inversion tropopause



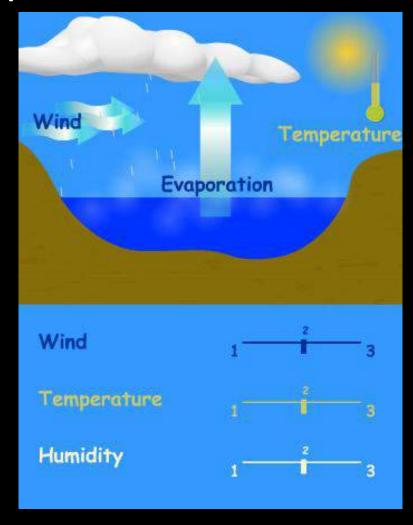
Temperature Inversion • Ex. Valley inversion





Precipitation process

Evaporation



Humidity

- Relative Humidity
- Absolute Humidity
- Specific Humidity

Precipitation process

Condensation- clouds



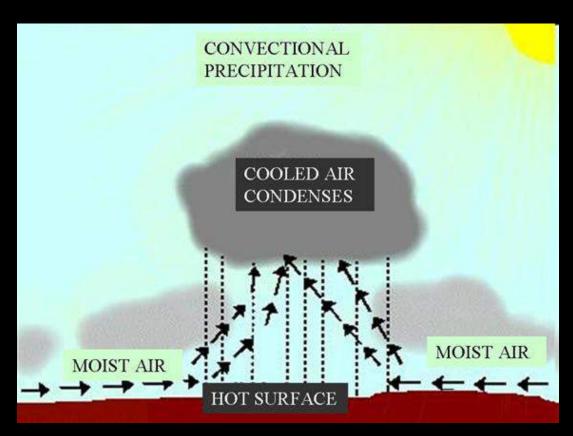
Type of Cloud

- High cloud
 Cirrus
 Cirrocumulus
 Cirrostratus
- 2. Medium cloud
 Altostratus
 Altocumulus
 Nimbostratus
- 3. Low cloud
 Stratus
 Cumulus
 Cumulonimbus

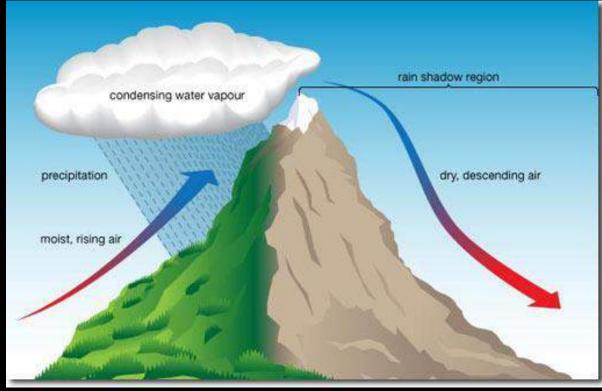
Discuss

Types of rainfall

Convectional rainfall



Orographic rainfall

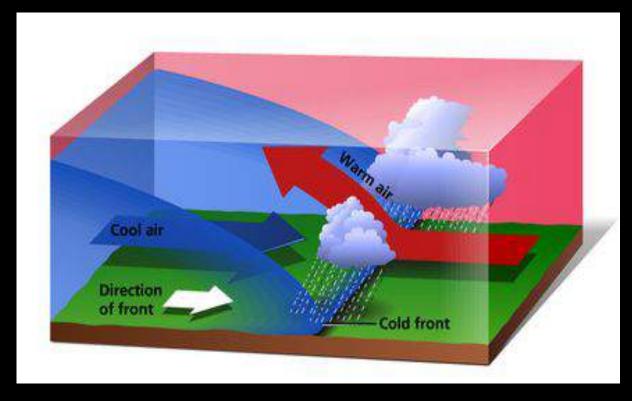


Types of rainfall

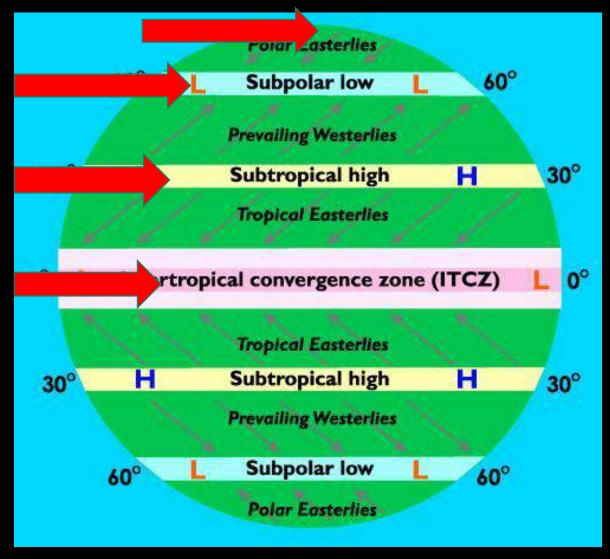
Cyclonic rainfall

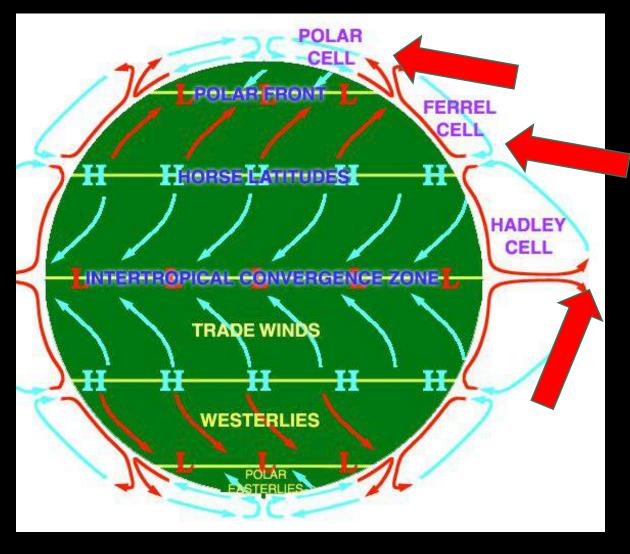


Frontal rainfall

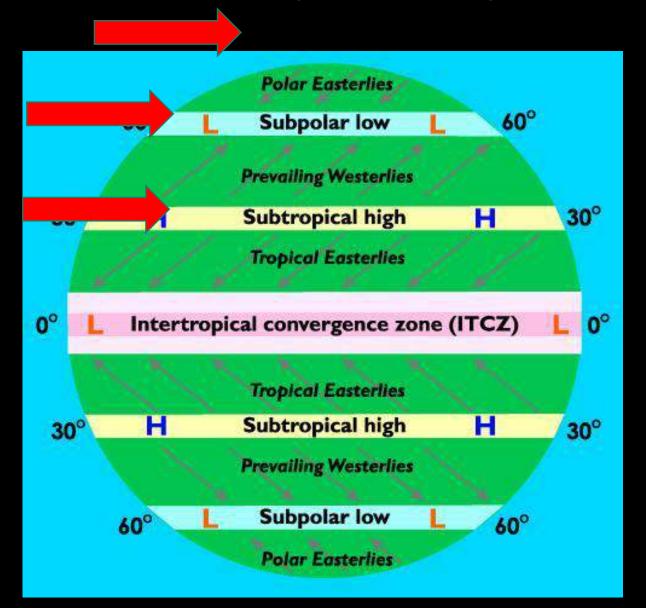


Pressure system of the world



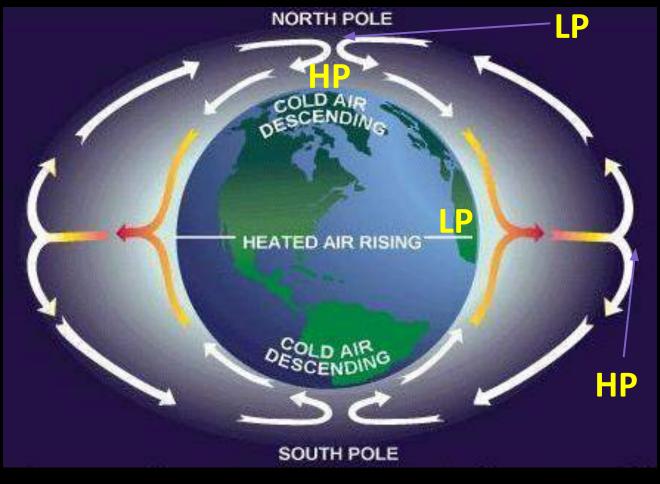


Planetary wind system



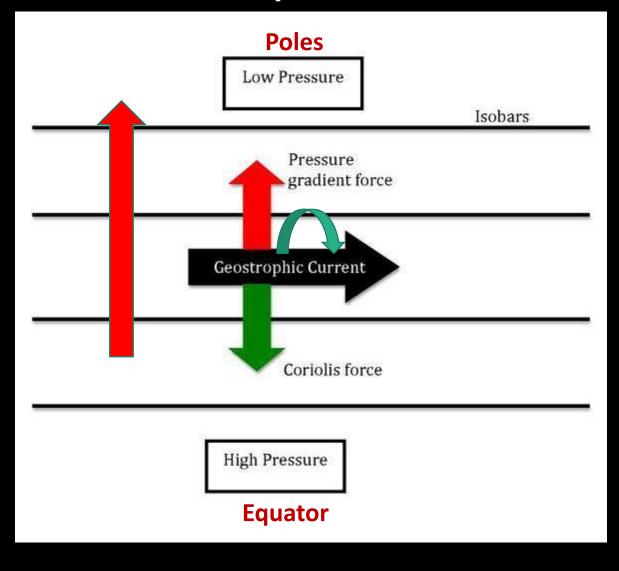
- High pressure to low pressure
- Coriolis force clockwise in northern hemi and anti-clockwise in southern hemi
- Trade winds
- Westerlies
- Polar easterlies

Upper tropospheric winds



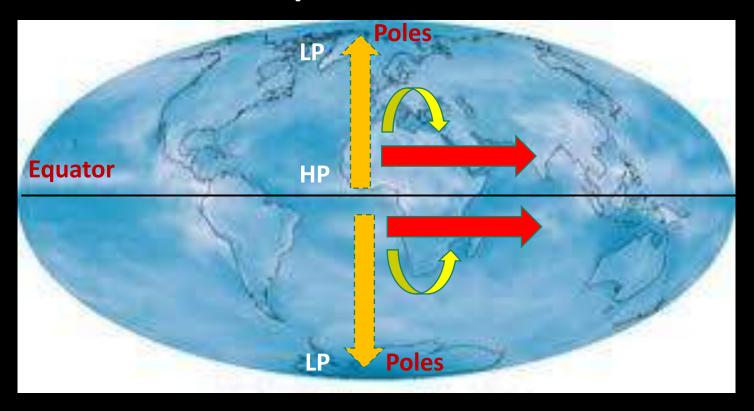
- Around tropopause, there is only one gradient
- Wind accumulated above equator and rarified atmosphere above poles
- HP at the equator and LP at the poles

Geo-strophic winds



- strong coriolis force at tropopause
- Because friction is less high speed - stronger the coriolis force
- So the deflection is 90 degree
- Such winds called geo-strophic winds

Geo-strophic winds



Northern hemisphere
 Geostrophic deflect clockwise
 west to east
 In southern Hemisphere:

Anti-clockwise movement

Deflect west to east

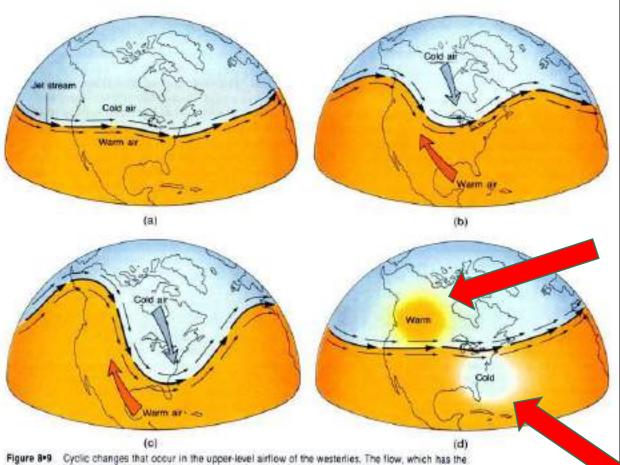
Westerlies winds

 The upper tropospheric winds / geo-strophic winds blow from west to east at the very high speed



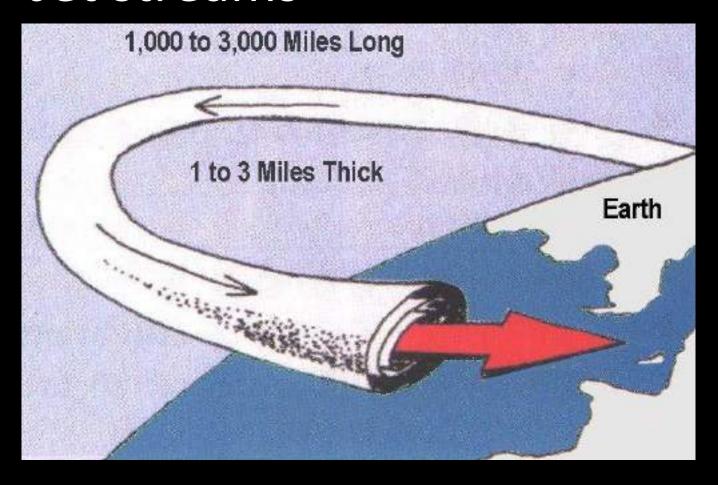
Rossby waves

jet stream as its axis, starts out nearly straight and then develops meanders that are eventually cut off, who



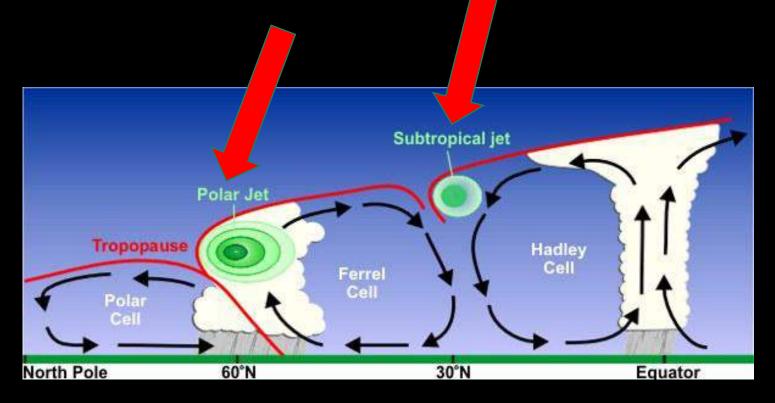
- Westerlies at poles to maintain the angular momentum- they meander => Rossby waves
- Rossby waves do not meander consistently, but follow a cycle = Index cycle

Jet streams



- In westerlies, there are strong, narrow bands of high speed wind => Jet stream
- Speed of Jet stream300kmph

Jet streams location



- There are situated at the margins of meridional cells
- 4 permanent Jet streams: 2
 Polar Jet and 2 Sub-Tropical
 Westerly Jet STWJ

Polar High **SPLP**

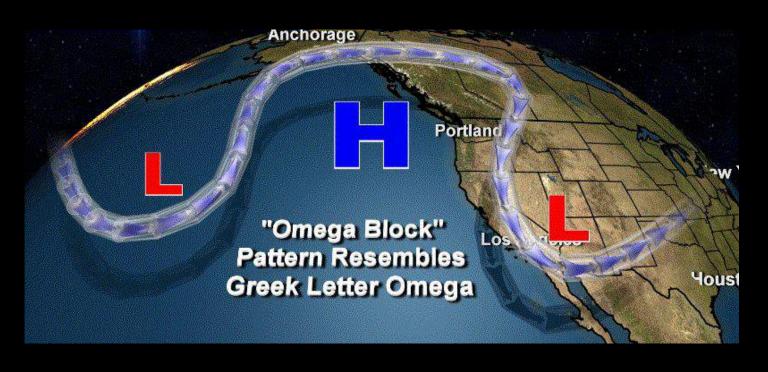
STHP

Eq. LP

Jet streams

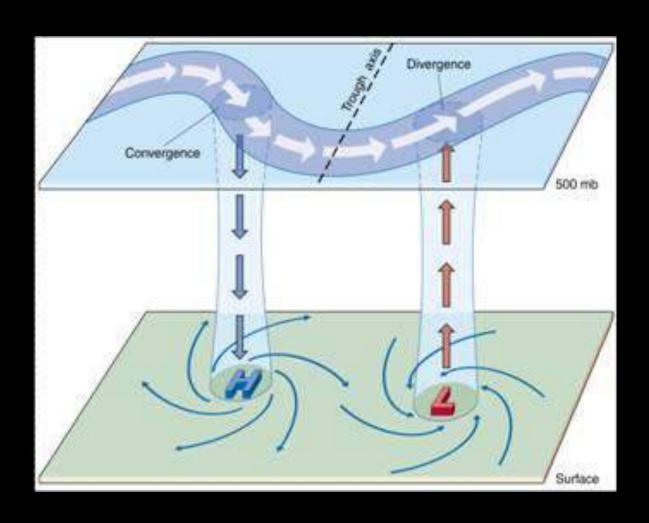


Jet streams



- Jet stream embedded in westerlies (Rossby waves) at high latitude, cause pressure variability
- That's why they are called travelling depression

Jet Stream – travelling depressions

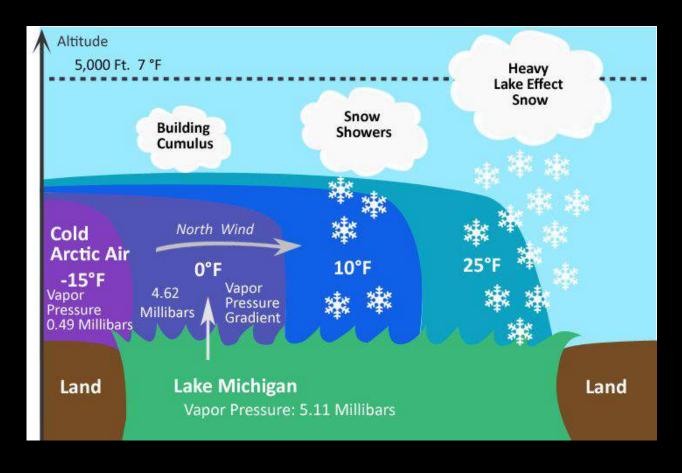


Any Enquire

Jet stream and Indian Monsoon

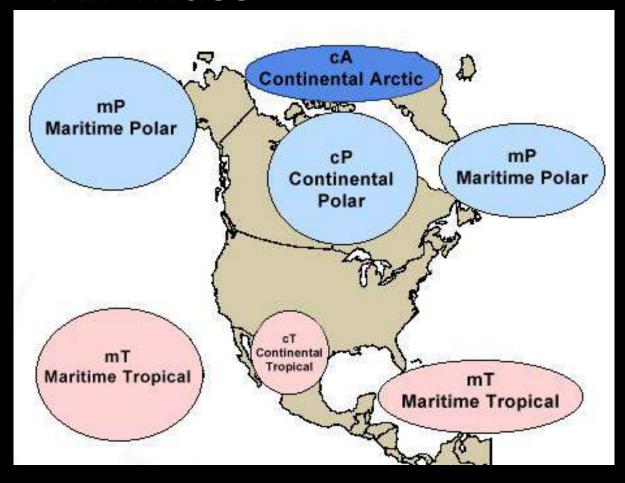
Break Time

Air mass



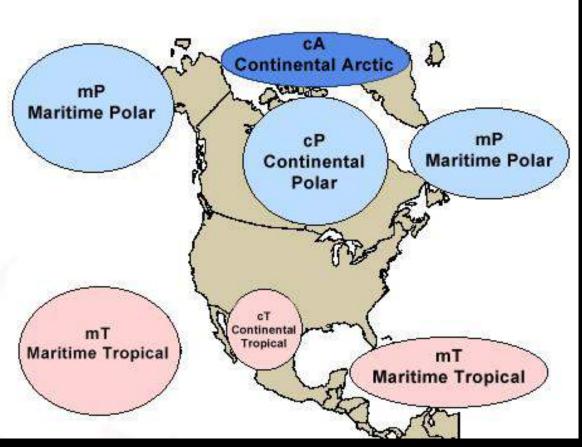
- Large extensive body of air-mass (1000sqkm)
- Height upto Tropopause
- At particular height, one air mass will have uniform temperature and moisture across its width
- Airmasses can be differentiate according to their temperature and moisture content

Air mass



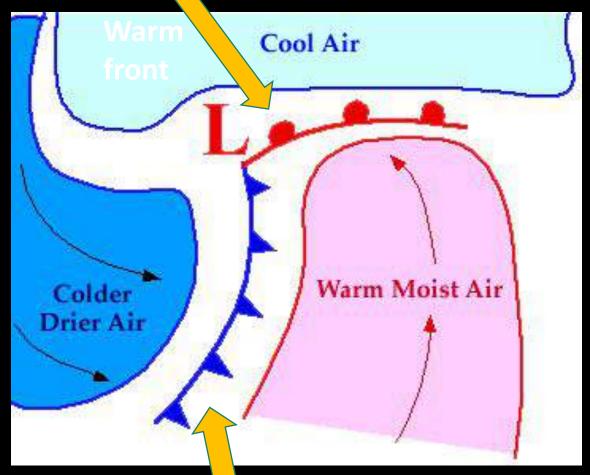
- Air mass acquired properties from the source regions – land, marine, polar, arctic, Antarctic = give them identity. Ex. mP, cT
- Extensive homogeneous surface
 + longer stay (HP)

Air masses



• Air masses do not stay at their source regions forever, they move out. While moving they came across other air masses.

Front

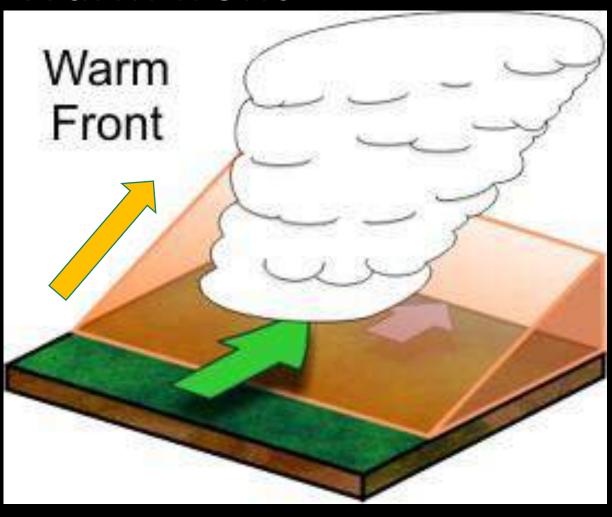


- The relative difference between temperature and moisture decide their interaction with one another
- The border/ meeting region of the two air-mass => Front

Cold front Cold Front

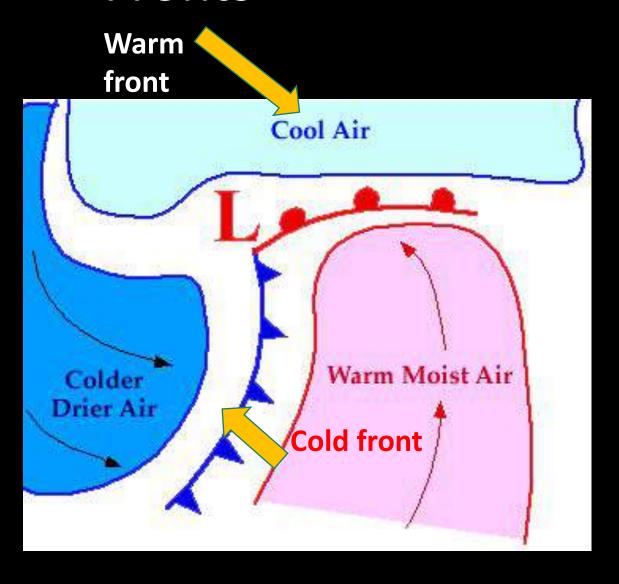
- If cold air mass move faster than the other than it will lift the warmer one upward => cold front
- the slope will be steep = there will be sudden up-liftment of the warm air = cumulonimbus clouds = cyclonic rain

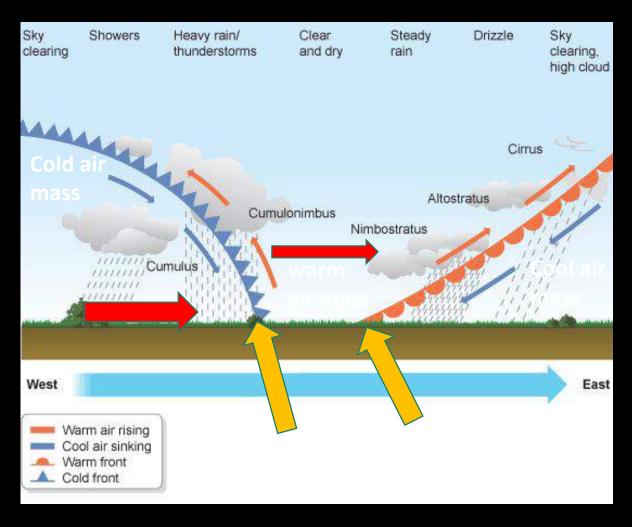
Warm front



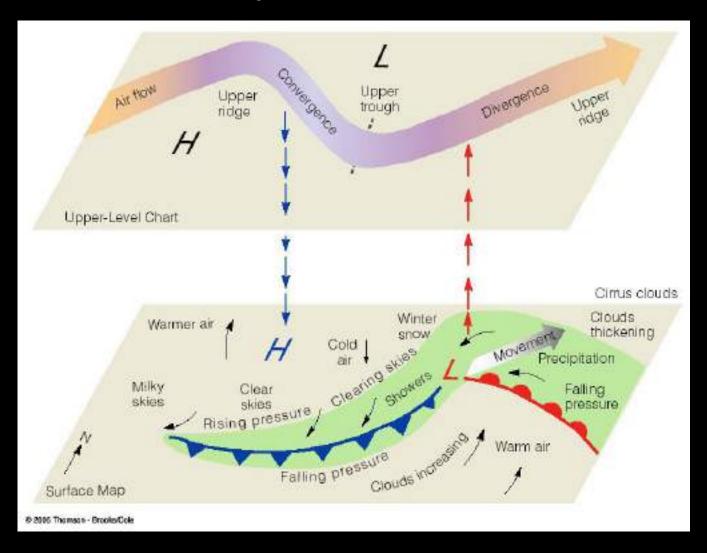
- If warmer air mass is more active than cold front => warm front
- slope will be gentler =
 there won't be sudden
 up-liftment of warm air =
 uniform prolonged rain –
 drizzle

Fronts





Frontal cyclone



 Also called as extra-tropical cyclone, travelling depressions, cold-core cyclone, wave cyclones

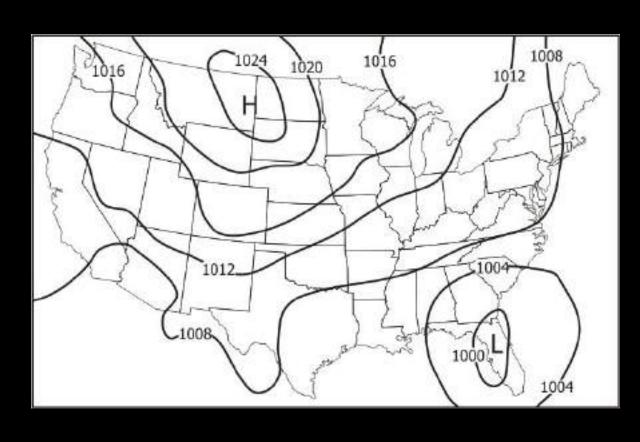
Meaning of cyclone

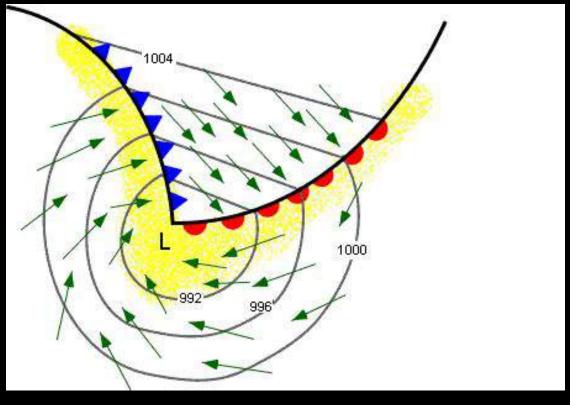


- 1) Intense LP system
- 2) Air converges towards the centre
- 3) In Northern hemisphere convergence anti-clockwise
- 4) Closed isobars

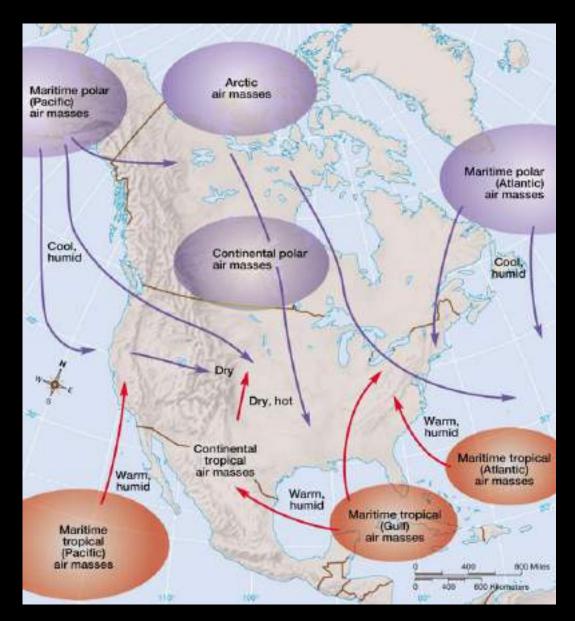
Isobar

Closed isobar



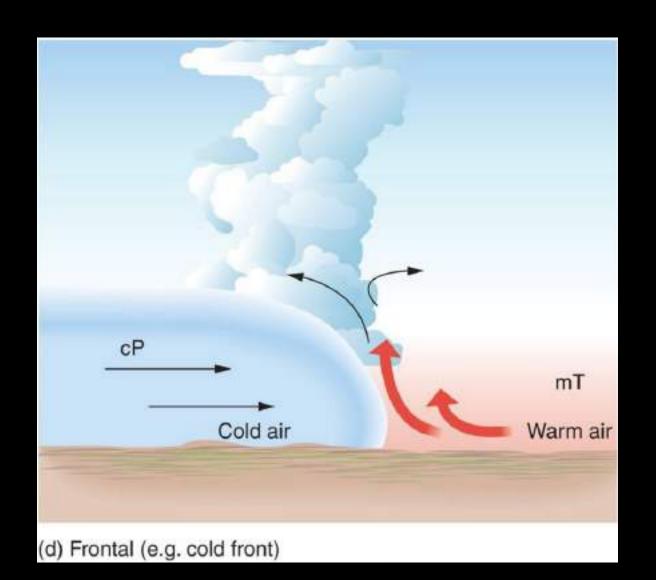


Development of Frontal cyclone



- Movement of airmasses from their source region
- The warm and cold air mass face each other
- A front is created between them
- Called Stationary front

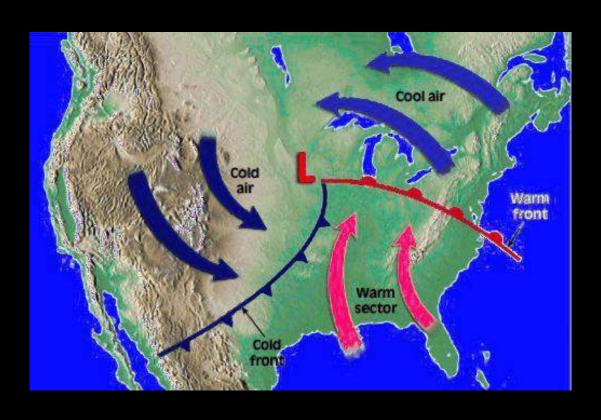
Development of Frontal cyclone

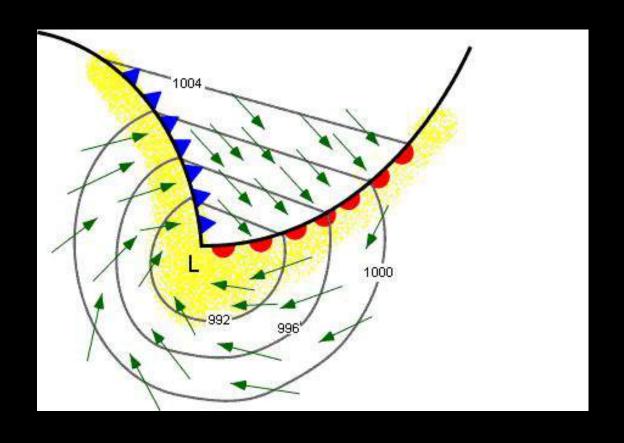


- Cold air mass pushed the warm air mass
- Forced upliftment of warm air mass at the cold front =LP
- Two cold air mass convergence circular due to coriolis force

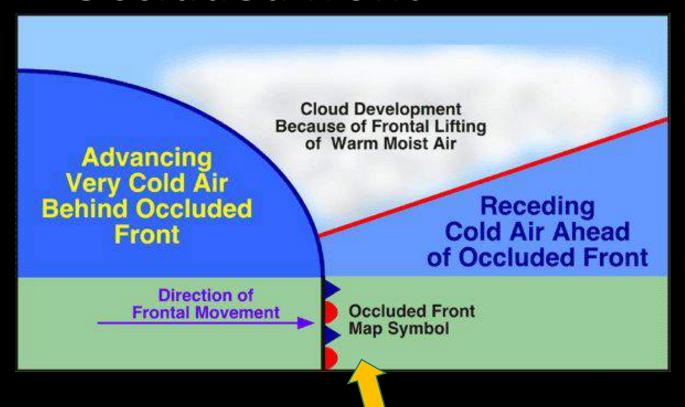
Mature stage

Interaction of air masses



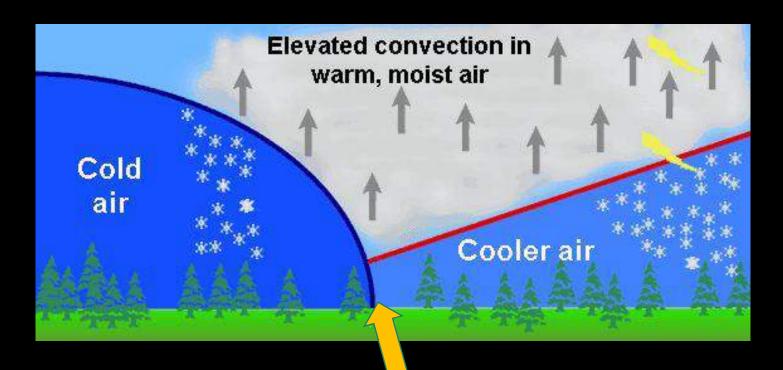


Occluded front



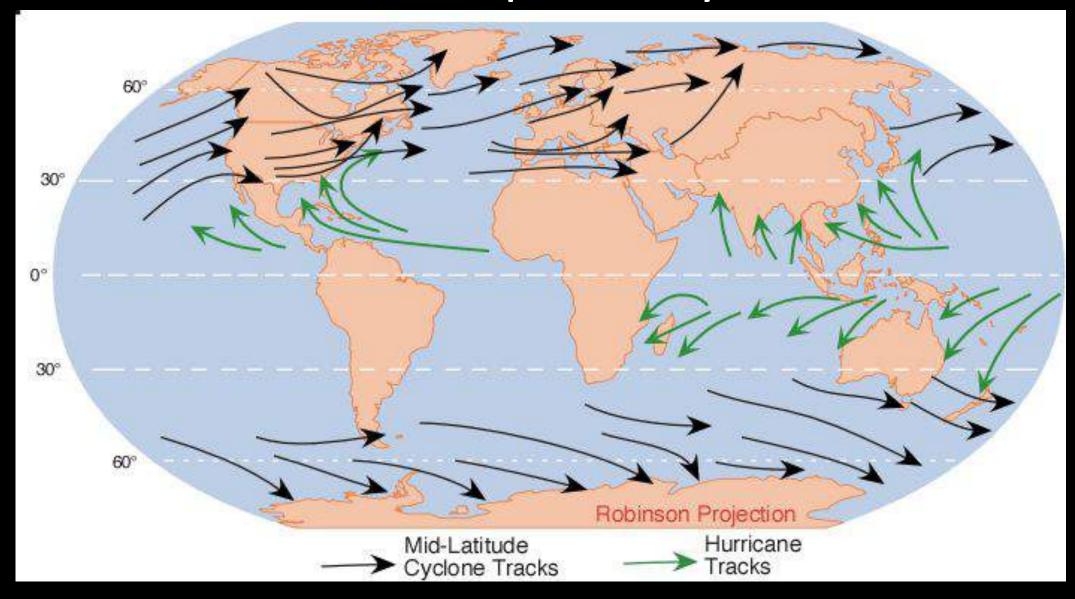
- One cold air mass climb over other cold air mass—warm front is destroyed
- Called occluded front
- Rapid change in temperature and pressure
- Unstable weather conditions

Dissipation of frontal cyclone . Frontolysis – no great

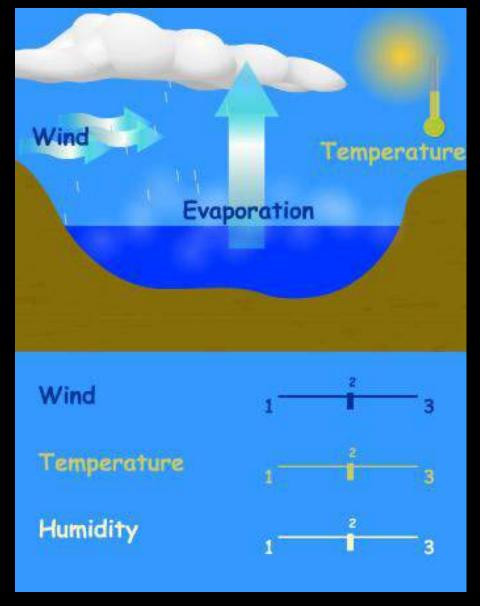


Frontolysis – no great temperature difference between two cold air masses – front dissipated – LP reduced – cyclone dissipated

Distribution of temperate cyclones



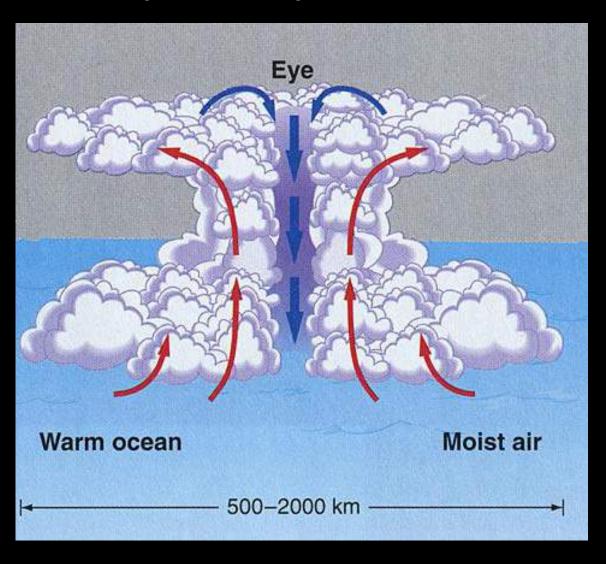
Tropical cyclone



Hurricane –N. America Typhoon - China

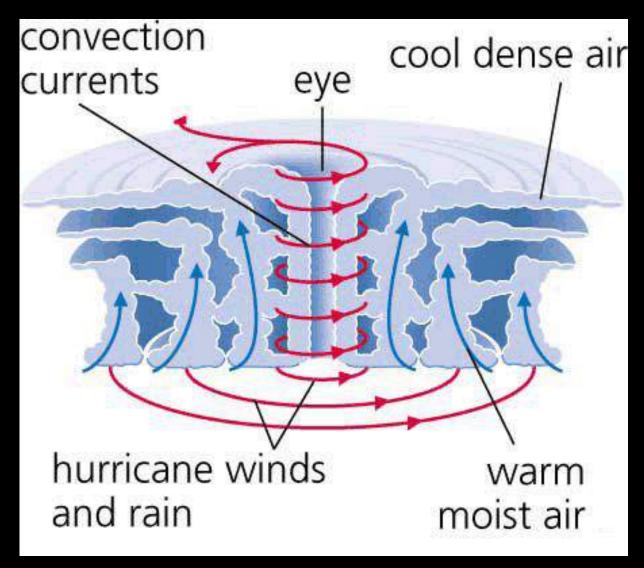
- Late summer
- •Increased sea surface temperature = LP

Tropical cyclone



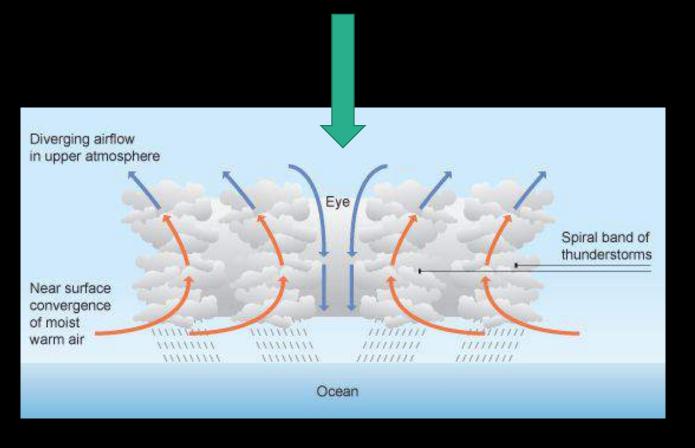
- Convergence of air around LP zone
- Rising moist (wet) air => absolute instability
- Cloud formation =more and more moisture – latent heat of evaporation => cumulo nimbus cloud => cyclone

Mature Tropical cyclone



- Intensification of LP
- Converging air near water surface
- Circulating air rises above (coriolis force)
- Diverging air at the top of cyclone

Eye of the tropical cyclone



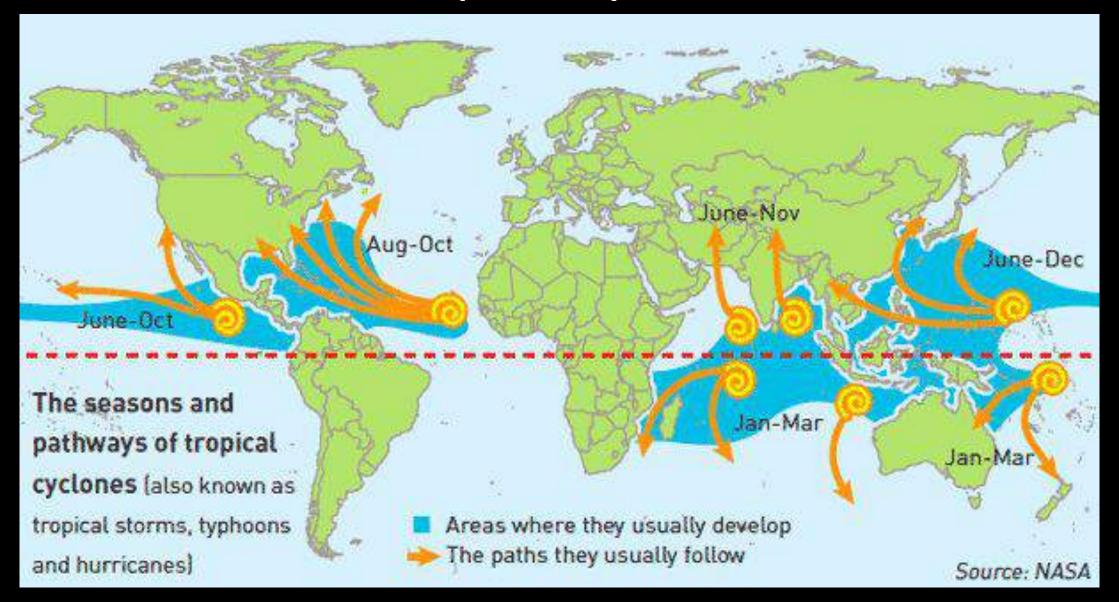
- At the centre of the cyclone 'eye' of the tropical cyclone.
- It is a pressure defect.
 Because, at 'eye' a narrow stream of wind descend = is
 HP at 'eye'
- At the eye, there is clear sky.
- Beyond eye wall extreme low pressure

Properties of tropical cyclones



- Move swiftly
- Always east to west
- It is fuelled by moisture so when cyclone is cut-off from sea and move towards land – it starts weakening

Distribution of tropical cyclone



comparison

Temperate cyclone

- 30-40 degree latitude
- Dynamically induced
- Due to frontal interaction
- Formed over large area
- Move west to east
- Gradual movement predictable

- Tropical cyclone
- 8-20 degree latitude
- Thermally induced
- Due to increasing SST
- Small area
- Move east to west
- Swift movement- difficult to predict path

comparison

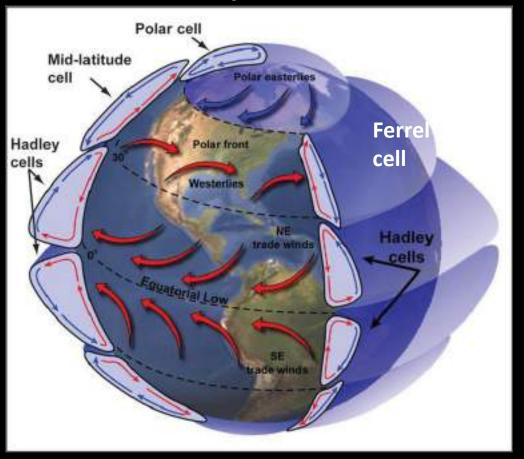
Temperate cyclone

- Wind speed 40-50 kmph
- Pressure gradient 980 mb
- Powerful on land
- Affect mainland
- More time to dissipate

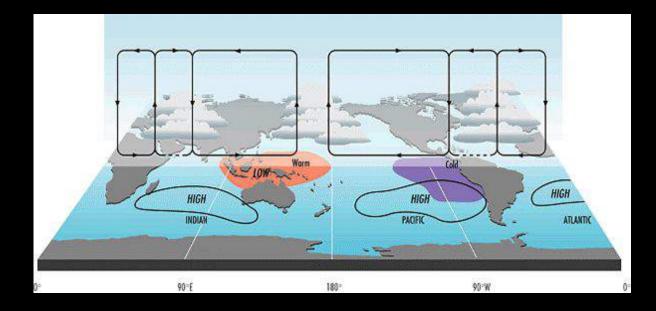
- Tropical cyclone
- Wind speed >120 kmph
- Pressure gradient <880 mb
- Weakens on land
- Affect only coastal areas
- Quickly dissipate after coming on land

Pressure gradient cells

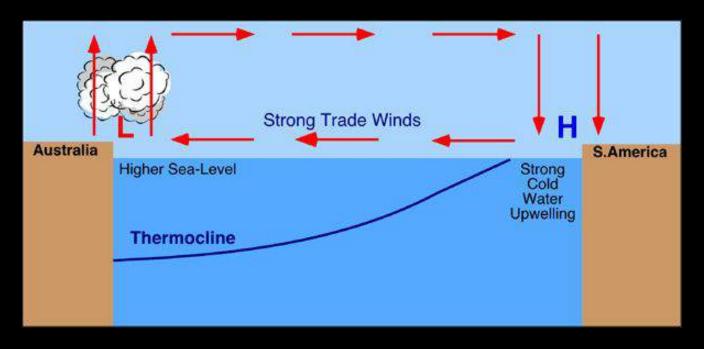
Vertical: Hadley, Ferrell



Horizontal: walker

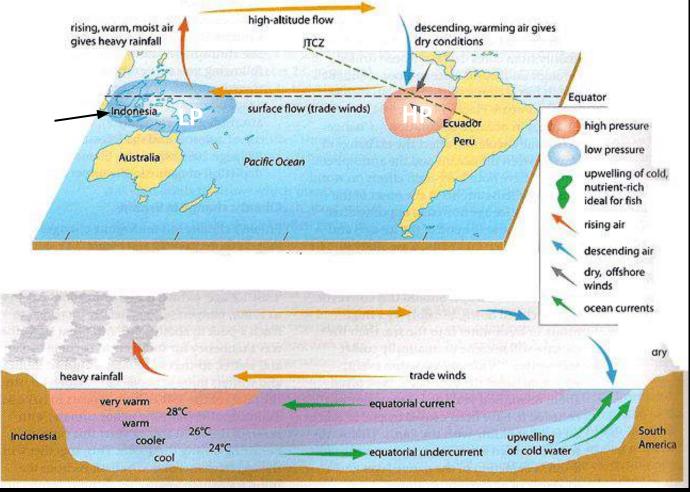


Walkar cell at south Pacific



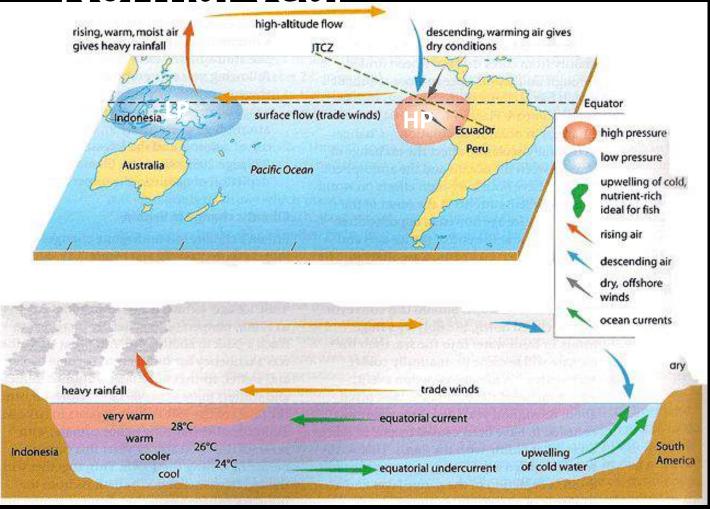
- There are oscillation in pressure gradient and air circulation after intervals of 2-3 years in south Pacific Ocean
- Called Southern Oscillation
- ENSO = EL Nino Southern Oscillation

Normal year



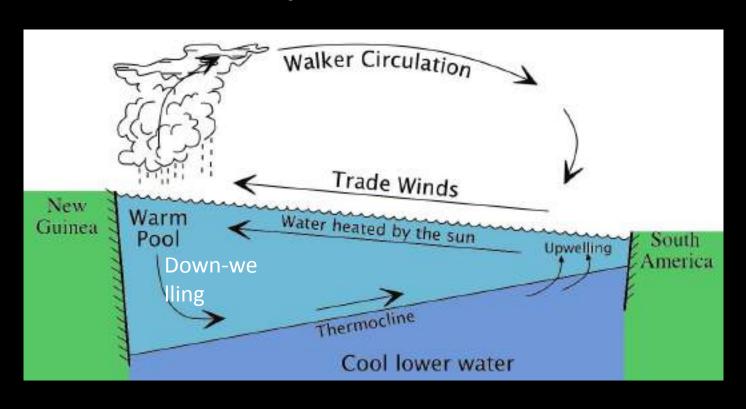
- South equatorial current pile up water at northern Australia – increase SST – called <u>West Pacific Pool</u>
- It brings rainfall in Northern Australia

Normal vear



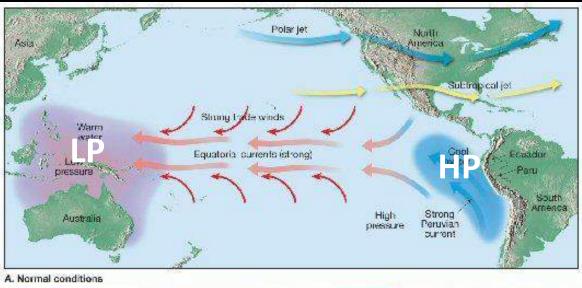
- The diverging air above Australia move towards Peruvian coast
- They descend at Peruvian coast = HP – desiccating effect to Atacama desert
- Completes the Walker cell

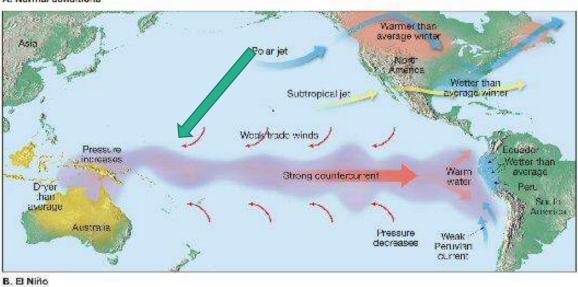
Normal year



- As south equatorial current take water from east to west, it led water from bottom to come up and take the space
- Up-welling at the Peruvian coast = rich fishing ground

El-Nino year

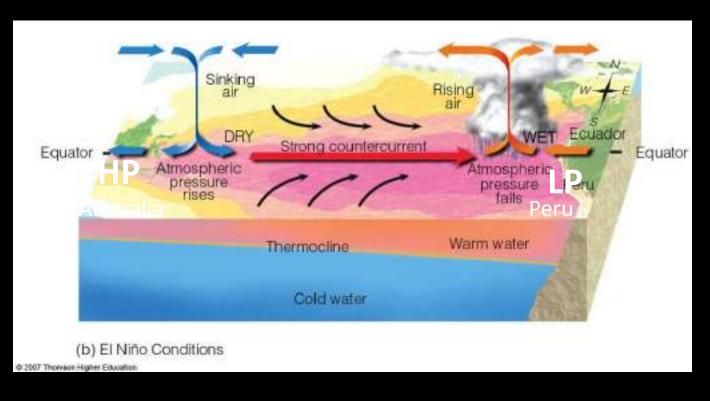




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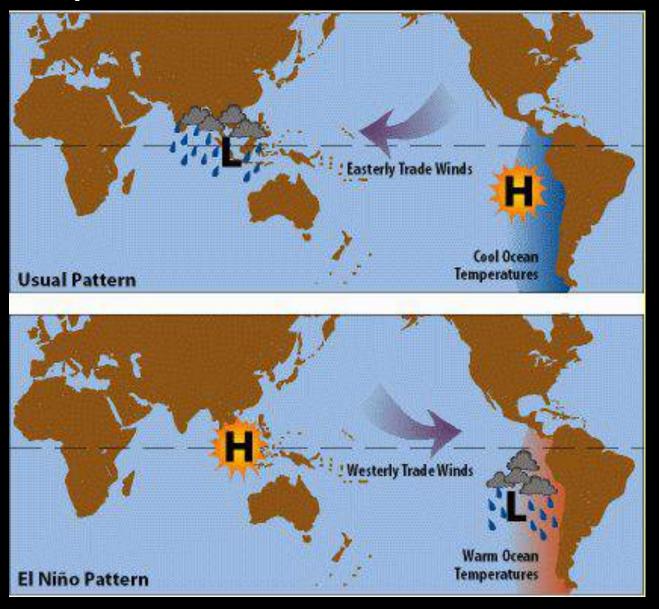
- Direction of walker cell reverses
- South equatorial current weakens (reason unknown)
- Weak piling up of water at Northern Australia
- Weakening of west Pacific Pool

El-Nino year



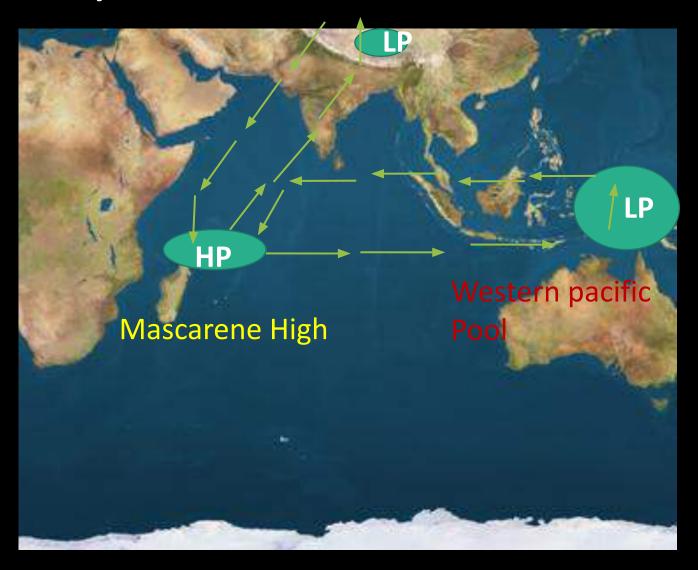
- Ocean water move towards
 Peruvian coast. Create LP system over there and rainfall at Atacama desert
- The rising and diverging wind above Peru descends over Australia = HP condition – drought in Northern Australia

Implications of El-Nino



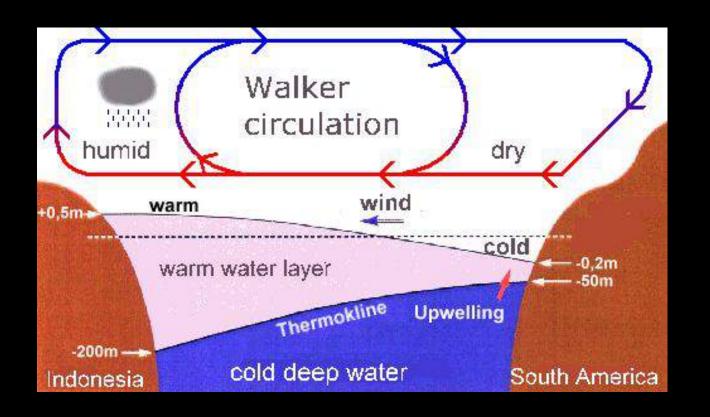
- El-Nino bring drought condition in Indonesia as well – forest fire
- It is responsible for weak monsoon in India

Implications of El-Nino on Indian Monsoon



- Association of El- Nino or Western Pacific Pool with Indian Monsoon
- There can be other atmospheric cells associated with Western Pacific pool
- Still undiscovered

La-Nina



- Intensification of walker cell
- Strong west pacific pool
- Heavy rainfall flood condition in Northern Australia – good monsoon in India
- Drought in Atacama
- Very good fishing business at Peruvian coast – price crash

Thank you